

# Supporting Information

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## SI Raman Spectroscopy

Raman spectra of organic matter were collected using a Kaiser HoloLab Series 5000 equipped with a diode-pumped solid-state laser (785 nm) (Utrecht University). Hyperspectral mapping was performed with a Horiba Scientific LabRam HR800 (Steinmann Institute, University of Bonn). Raman scattering was excited with a diode-pumped solid-state laser (784 nm) with less than 40 mW at the sample surface. A 100× objective with a numerical aperture of 0.9 was used resulting in a diffraction-limited lateral resolution of  $\sim 1 \mu\text{m}$ . The confocal hole was set to 1,000  $\mu\text{m}$ , resulting in a depth resolution of a few micrometers. The scattered Raman light was collected in a 180° backscattering geometry by an electron-multiplier charged coupled device detector after passing through a 100- $\mu\text{m}$  (single measurements) or 200- $\mu\text{m}$  (mapping) spectrometer entrance slit and being dispersed by a grating of 600 grooves per millimeter, yielding a spectral resolution of 1.7 and 2.3  $\text{cm}^{-1}$ , respectively, as determined from the width of Ne lines. The spectrometer was calibrated with the first-order Si Raman band at 520.7  $\text{cm}^{-1}$  and Ne lines. The total acquisition time varied between 900 and 1,200 s for single measurements and 25  $\times$  0.5 s for mapping. Normalization of Raman spectrum to the most intense epoxy band at 815  $\text{cm}^{-1}$  shows no contribution of bands from the epoxy resin (Fig. S1). This is evident by the absence of epoxy bands (arrow in Fig. S1) between 950 and 1,000  $\text{cm}^{-1}$  in the sample spectra. The false-color Raman image shown in Fig. 2G is generated from the spectra recorded for each pixel of the image by color-coding the ratio between the integrated intensity of an organic vibrational modes near 639  $\text{cm}^{-1}$  and a serpentine mode at 690  $\text{cm}^{-1}$  (Fig. S2). Warm colors reflect a high content of organic material within the analyzed volume, whereas blue colors mark regions with low or undetectable organic material. Spectra were collected over an area of 60  $\times$  60  $\mu\text{m}^2$  with a step size of 0.5  $\mu\text{m}$ , resulting in 7,200 spectra. The resulting map is shown in Fig. S4. To obtain reduced or  $R(\omega)$  intensities for future comparison that are independent from the instrument, the laboratory temperature, and the excitation wavenumber,  $\nu_e$ , and thus directly proportional to relative scattering activity, the Raman spectra from the mesh core were also corrected for (i) the instrument response function (white light correction); (ii) the excitation frequency dependence, that is, by the scattering factor,  $(\nu_e - \nu)^3$ , with  $\nu$  being the wavenumber of the scattered light (intensities were measured in photons per second); and (iii) the temperature effect, that is, by the Bose-Einstein temperature factor,  $1 - \exp(-h\nu_c/kT)$ , with  $h$ ,  $c$ ,  $k$ ,

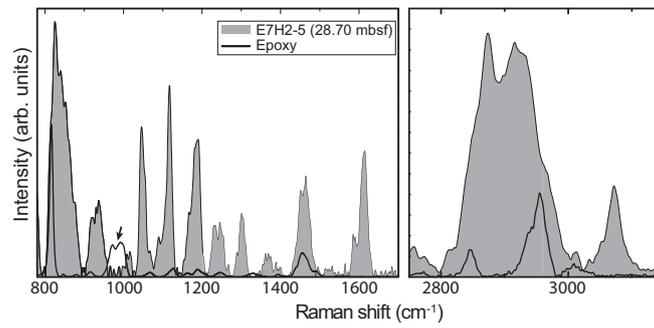
and  $T$  being the Planck constant, the speed of light, the Boltzmann constant, and the temperature, respectively. The corrected spectra were deconvoluted by least-squares fitting Gauss-Lorentz functions along with a linear background that was subtracted after the least-squares fit (Fig. S5 and Table S1). The error of the reported band positions is smaller than 0.1  $\text{cm}^{-1}$ . Raman evaluation of the serpentine polymorphs present in the rim and core structures was conducted by examining the water vibration region at  $\sim 3,600 \text{ cm}^{-1}$ . Raman spectra taken, using the 785-nm laser, showed no evidence for the OH bands due to the low scattering efficiency of these bands with this laser. Therefore, the 532-nm coherent compass sapphire laser of a WITec alpha 300R was used for the OH band analysis of serpentine group minerals. The analysis was conducted with a 50× long working distance lens with a 0.55 numerical aperture in backscattering geometry to the sample. After the scattered Raman light passed through a pinhole of 20  $\mu\text{m}$ , the light was dispersed on a grating of 1,800 grooves per millimeter, resulting in a spectral resolution of 1.1  $\text{cm}^{-1}$  in the spectral region of interest, measured using a built-in calibration light source.

## SI Opaque Mineral Grain Analysis

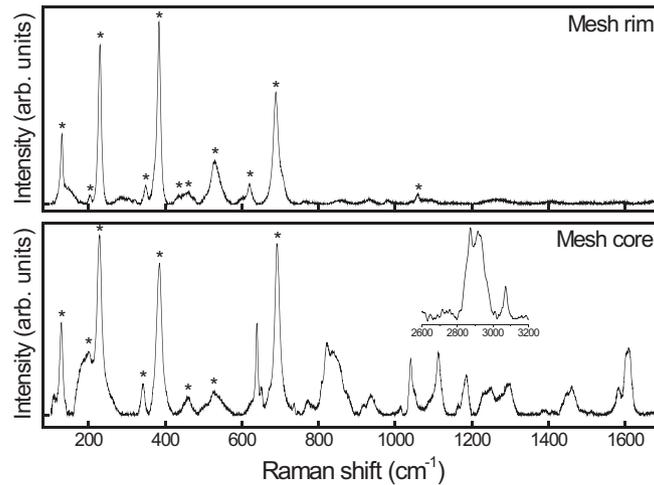
Fig. S6 shows the distribution of opaque mineral grains (high backscattering intensity in Fig. S6A) within the serpentinites clasts. Several FIB-SEM sections were prepared across the mesh rim-core interfaces, where Fig. S6B shows a representative high angle annular dark field (HAADF) image highlighting three nanoparticles with high intensity (overview image in main-text Fig. 3C). The intensity in an HAADF image scales with atomic number, implying that the nanoparticles are atomically heavier than the surrounding serpentine grains. Energy-dispersive X-ray analysis (Fig. S6C) reveals that the nanoparticles exist exclusively of Ni and Fe in a  $\sim 2:1$  ratio. Together with chemical analysis conducted on the micrometer-scale in a scanning electron microscope, the microstructural association of the particles within the serpentinites, and the limited possibilities of potential mineral phases with Ni and Fe in a  $\sim 2:1$  ratio (65), we conclude that the nanoparticles are awaruite ( $\text{Ni}_2\text{Fe}-\text{Ni}_3\text{Fe}$ ).

## SI Estimation of the Maximum Depth for the Current Temperature Limit for Life

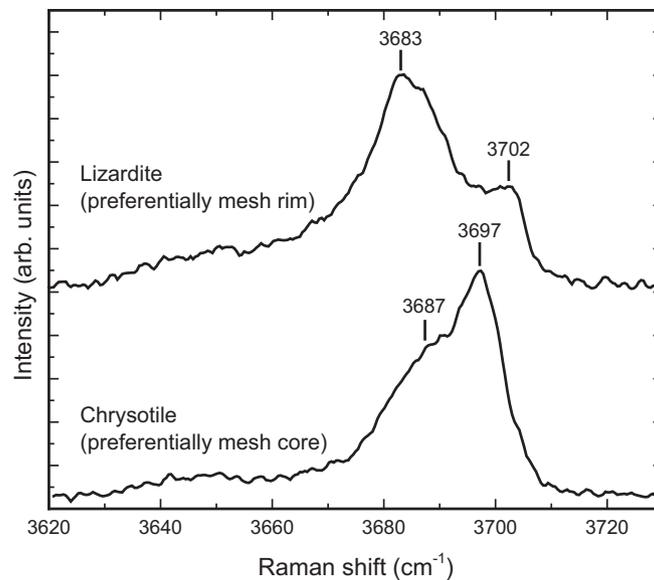
A sketch of the one-dimensional steady-state heat conduction model to estimate the maximum depth for the current temperature limit for life in subduction zone forearcs is found in Fig. S7.



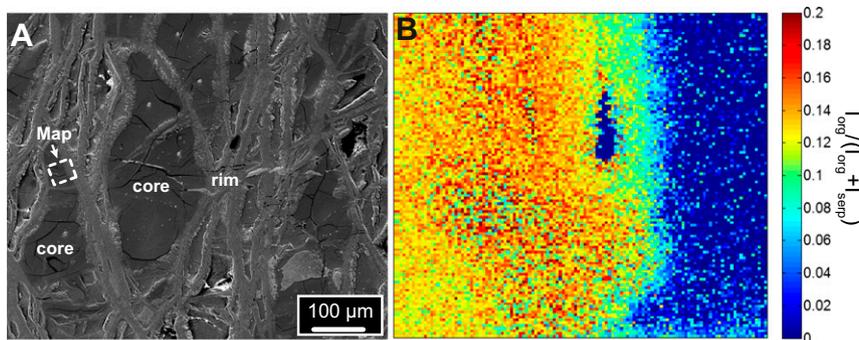
**Fig. S1.** Epoxy-normalized Raman spectrum of sample E7H2-5 (28.70 mbsf). All spectra were normalized to highest intensity epoxy Raman mode (815 cm<sup>-1</sup>), showing that epoxy does not contribute to the observed Raman spectra of the identified organic molecules.



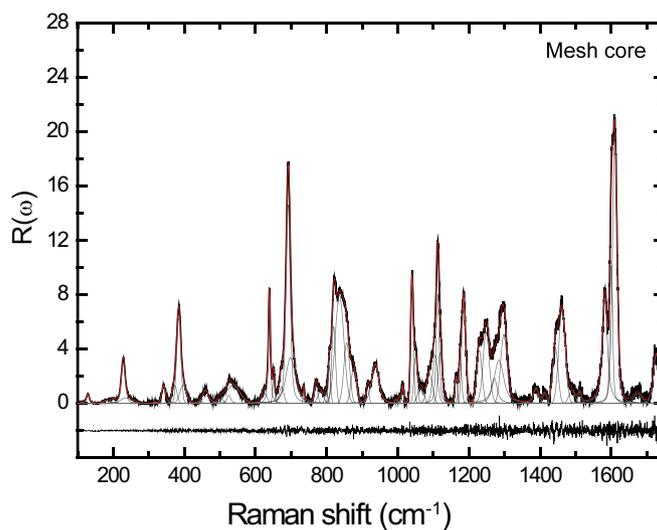
**Fig. S2.** Representative Raman spectra of the mesh core and rim region taken from the area in which hyperspectral imaging was performed. Bands marked with an asterisk belong to lizardite/chrysotile. All other bands reflect complex organic material (see main text and Table S1).



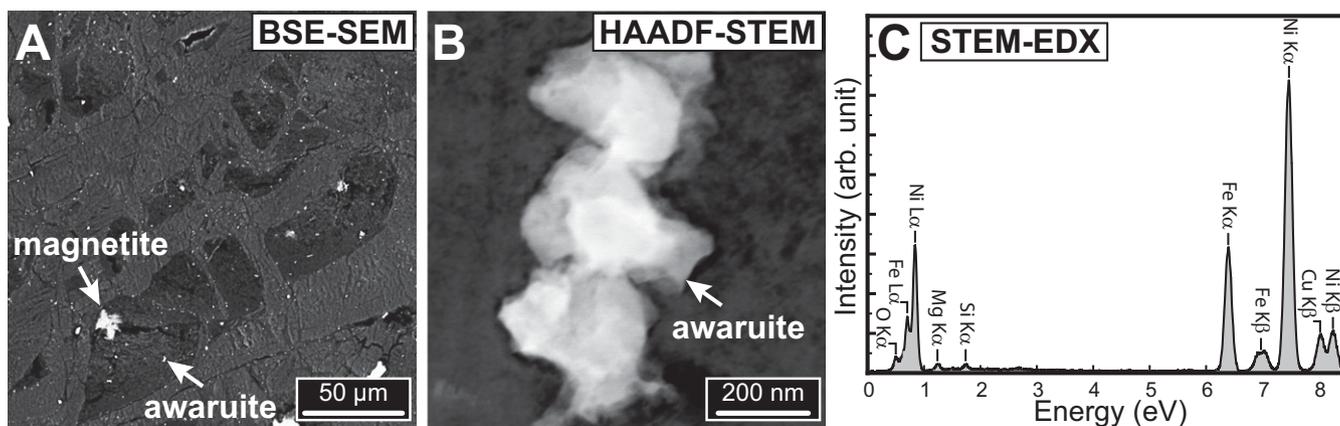
**Fig. S3.** Representative Raman spectra of OH-stretching modes, fingerprinting lizardite and chrysotile (e.g., ref. 64) typically found in the mesh rim and core regions, respectively, of the serpentinite clasts.



**Fig. 54.** (A) Backscattered electron image showing the location of the Raman map. (B) Shown is the distribution of organics within the mesh core and rim.  $I_{639}$  and  $I_{690}$  are the integrated intensities of the bands near  $639$  and  $690$   $\text{cm}^{-1}$ , respectively.



**Fig. 55.** Raman spectrum as shown in Fig. S2 from the mesh core, but with reduced intensities  $R(\omega)$  that were obtained by correcting the measured intensities for the instrumental response function, temperature effects, the excitation frequency dependence, and background (for more details, see *Methods* in the main text). Also shown is the deconvolution of the spectrum obtained from least-squares fitting individual Gauss-Lorentzian functions (gray curves) to the data. The red curve represents the sum curve. The residuals of the fitting procedure are also shown.



**Fig. 56.** A is a backscattered electron (BSE) image taken in a scanning electron microscope showing the distribution of opaque minerals (high backscattering intensity). B is a high-angle annular dark-field (HAADF) image taken with transmission electron microscope in scanning mode. The corresponding EDX analysis of a nanosized awaruite grain is shown in C. The Cu K $\beta$  peaks originates from the FIB section sample holder. The Mg K $\alpha$ , Si K $\alpha$ , and O K $\alpha$  peaks are a minor contribution from the surrounding serpentine grains.

